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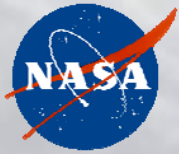
Summary of EDMF (and some PDF cloud parameterization) research at JPL and Caltech

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Summary of EDMF (and some PDF cloud parameterization) research at JPL and Caltech

Science Topics:

EDMF and dry convective boundary layer - Marcin Witek, Joao Teixeira

EDMF and moist shallow convection - Kay Suselj, Joao Teixeira

EDMF and PDF-clouds in a simplified GCM to study climate change -
Zhihong Tan, Remi Lam, Tapio Schneider, Joao Teixeira

LES research - Georgios Matheou, Daniel Chung



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Summary of EDMF (and some PDF cloud parameterization) research at JPL and Caltech

Projects and funding:

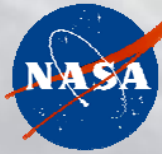
JPL LES model development and evaluation (JPL internal) - 2008-2011

EDMF development and implementation in US Navy mesoscale model
COAMPS (ONR) - 2008-2011

EDMF and PDF-clouds development and implementation in NASA GMAO
model (NASA) - 2009-2012

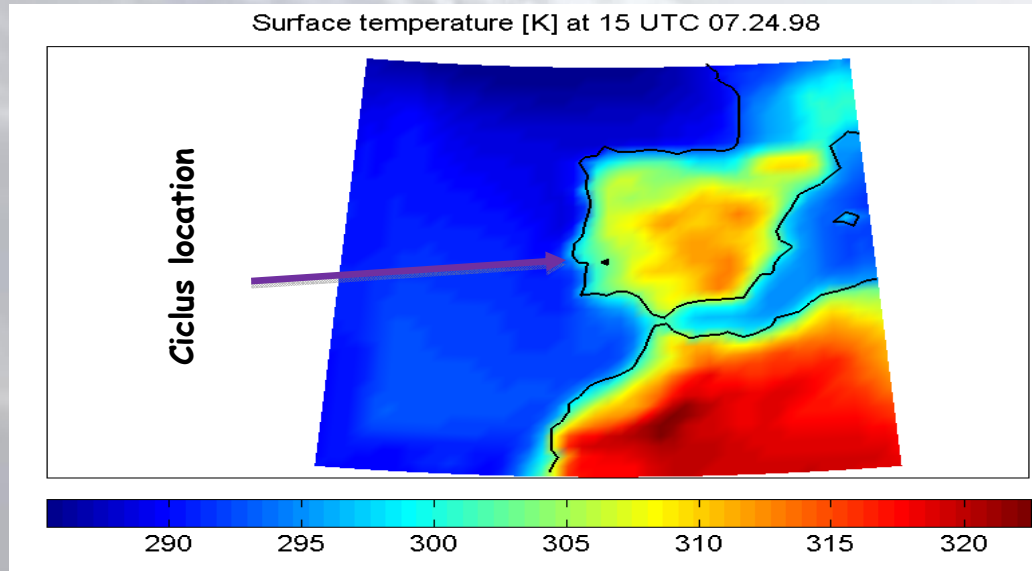
EDMF development and implementation in NCEP model (NOAA CPT) -
2010-2013

EDMF and PDF-clouds development and implementation in US Navy global
model NOGAPS (ONR) - 2011-2014 ³

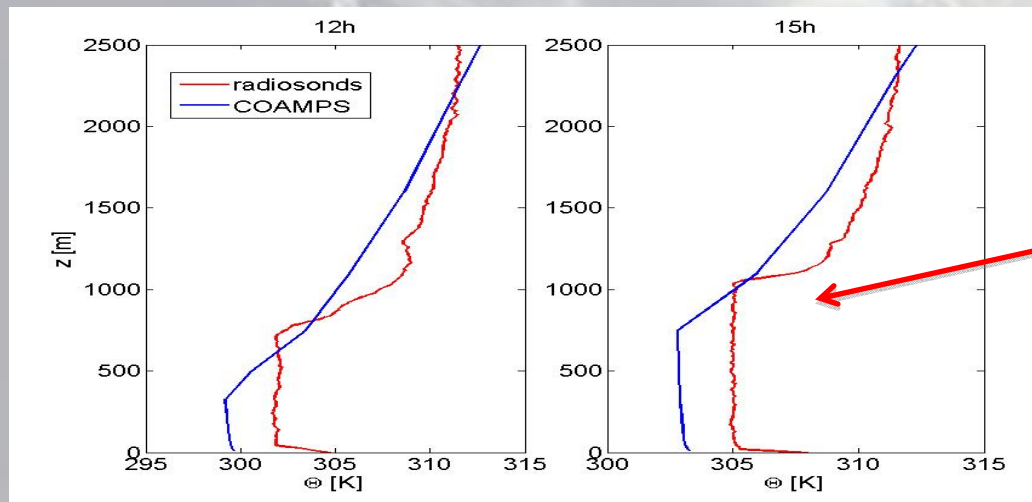


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Dry Convective Boundary Layers: example of US Navy mesoscale model COAMPS



Example - CICLUS experiment



COAMPS control:

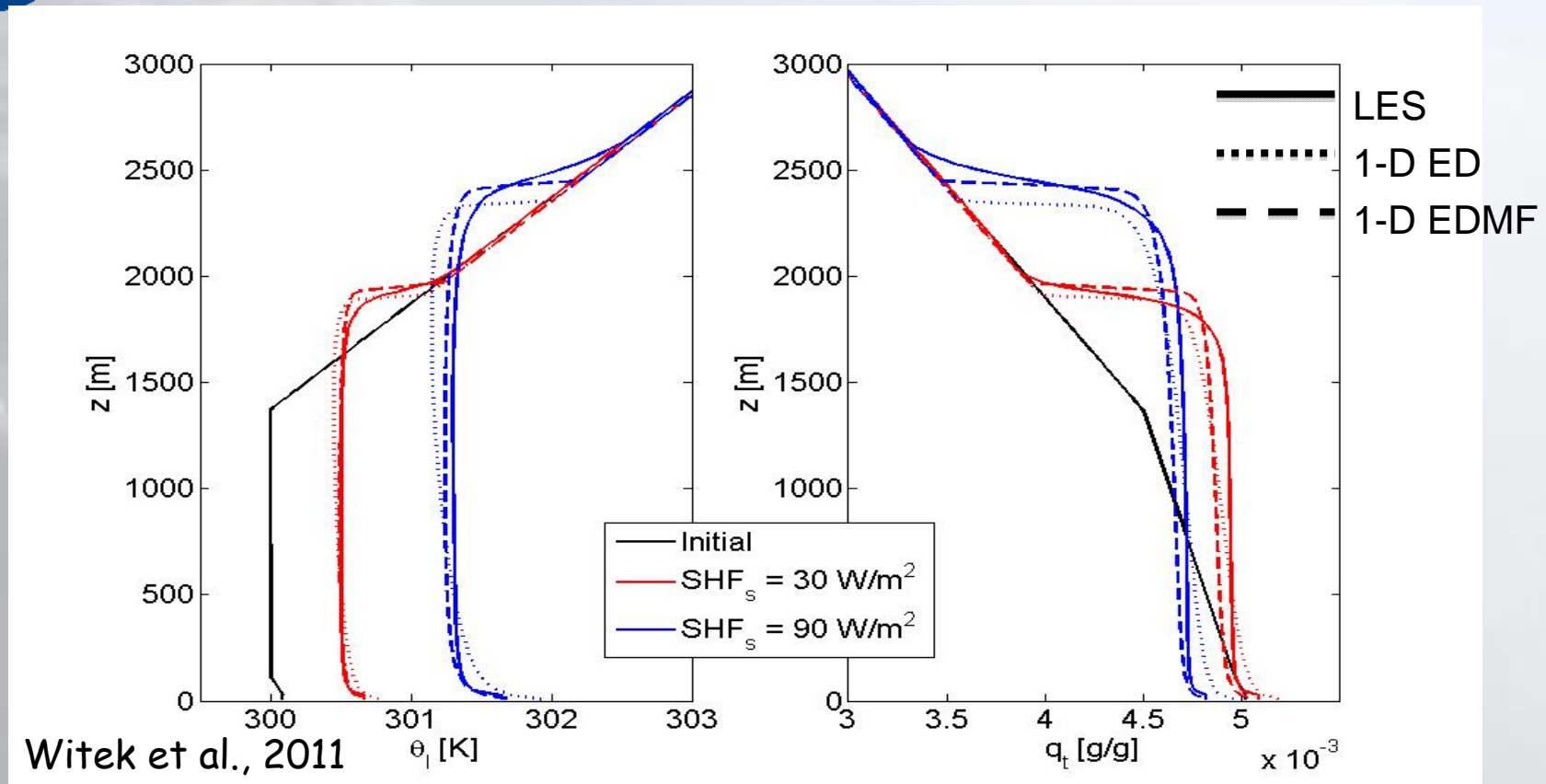
- not enough entrainment (too low and too cold boundary layer)
- profiles not well mixed

Teixeira et al., 2004



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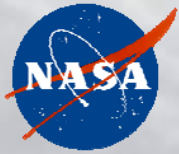
Dry Convective Boundary Layer: θ and q_t vertical profiles after 6 hours with EDMF and TKE



Full EDMF simulations:

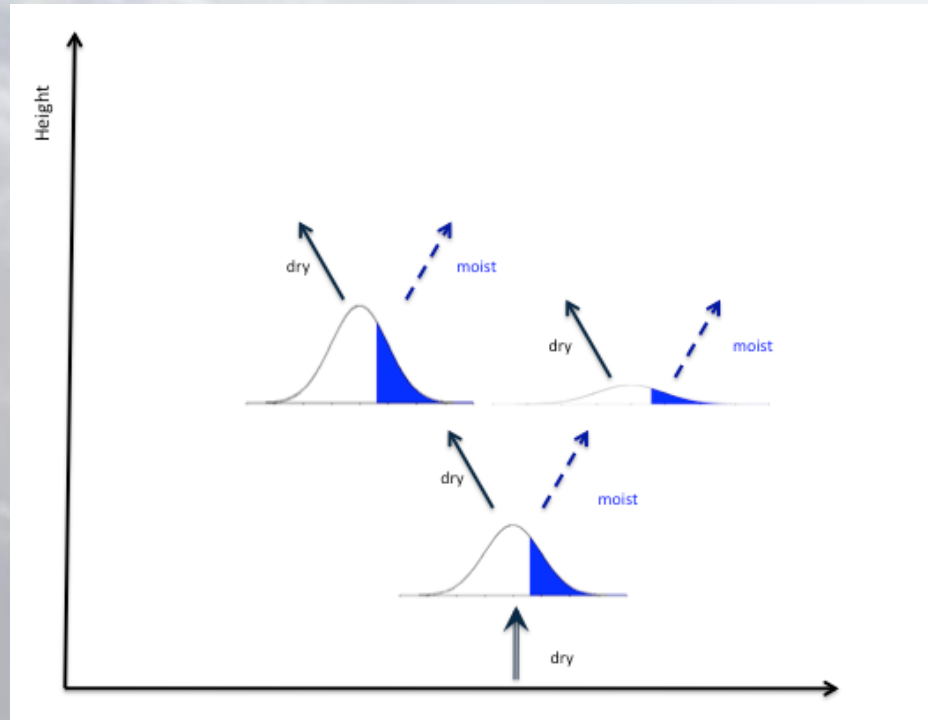
Witek et al, JAS, 2011

- surface layer more realistic
- neutral profile in the well-mixed layer
- larger entrainment leads to better inversion height
- inversion layer too sharp compared to LES



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EDMF and Shallow convection: using PDF of updraft properties



Variance of updraft PDF:
Local balance between
production and destruction

Partial updraft condensation:

- Start with single dry updraft at surface, integration in vertical
- Using updraft PDF to estimate updraft cover and water at each level
- Separation of dry and moist updraft when partial condensation occurs
- Moist and dry updraft-areas are integrated independently in vertical (with different vertical velocities)

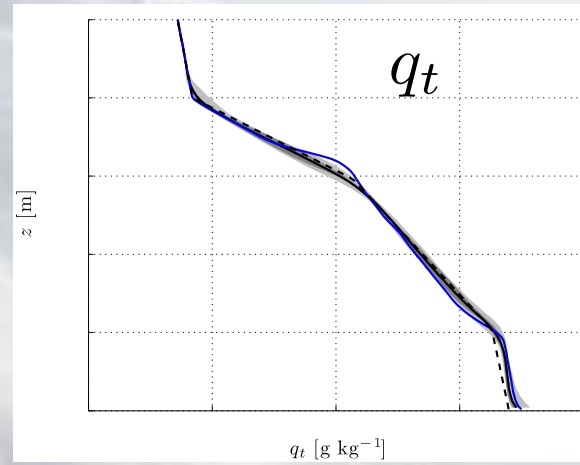
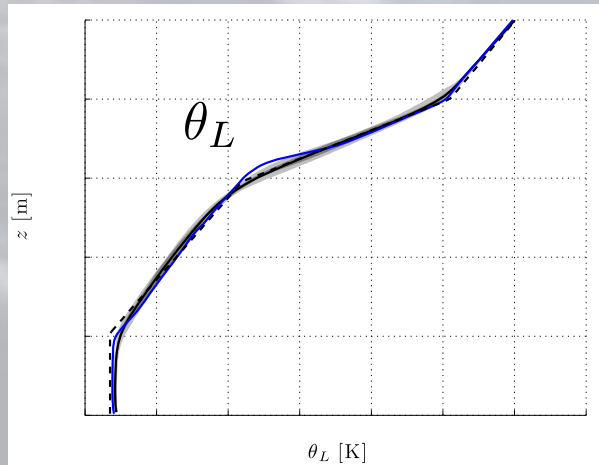
Provides estimation of updraft area and avoids need for cloud base closure



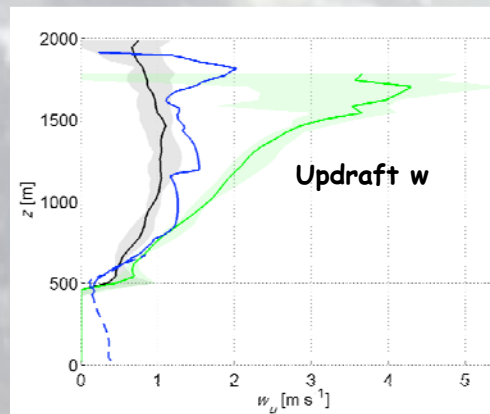
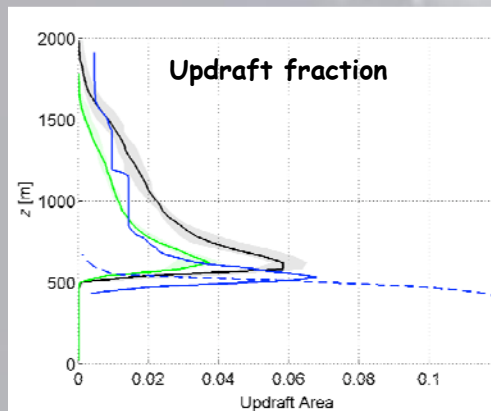
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EDMF simulation of shallow cumulus BOMEX case: comparison with LES

Mean profiles between 3rd and 4th simulation hour



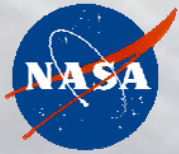
— Single column model
- - LES, mean



— Single column model, dry
— Single column model, moist
— LES, cloud core, mean
— LES, cloud core, range
— LES, clouds, mean
— LES, clouds, range

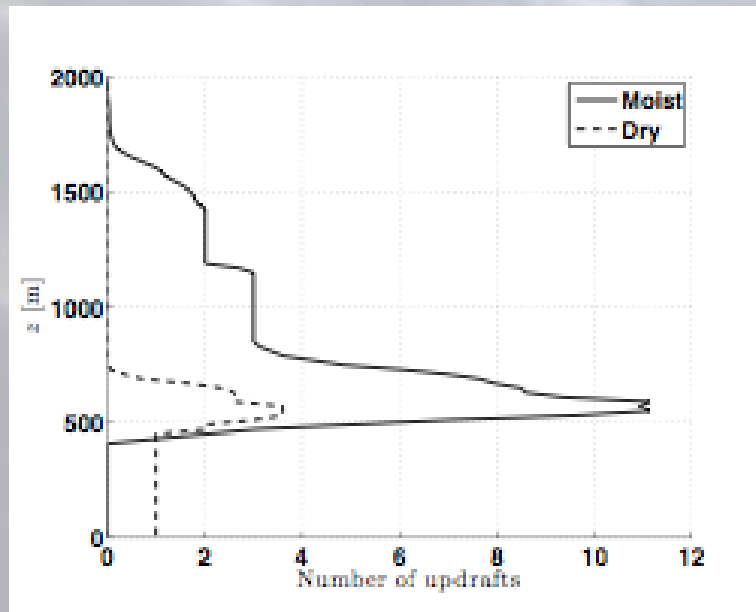
Suselj et al, JAS, 2011

New aspect: Using PDF of updraft properties

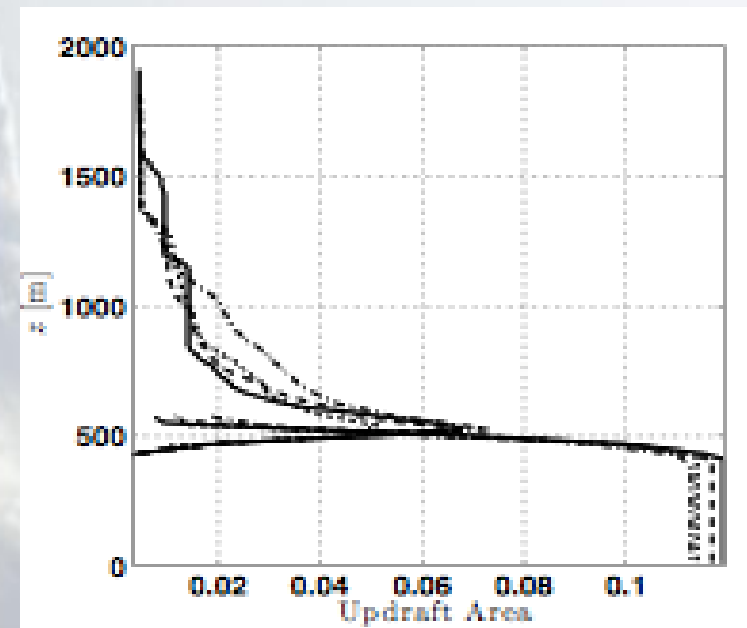


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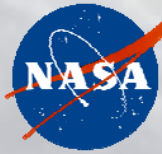
EDMF using PDF of updraft properties: BOMEX and the sensitivity to vertical resolution



Number of updrafts for control
simulation (DZ=20 m)

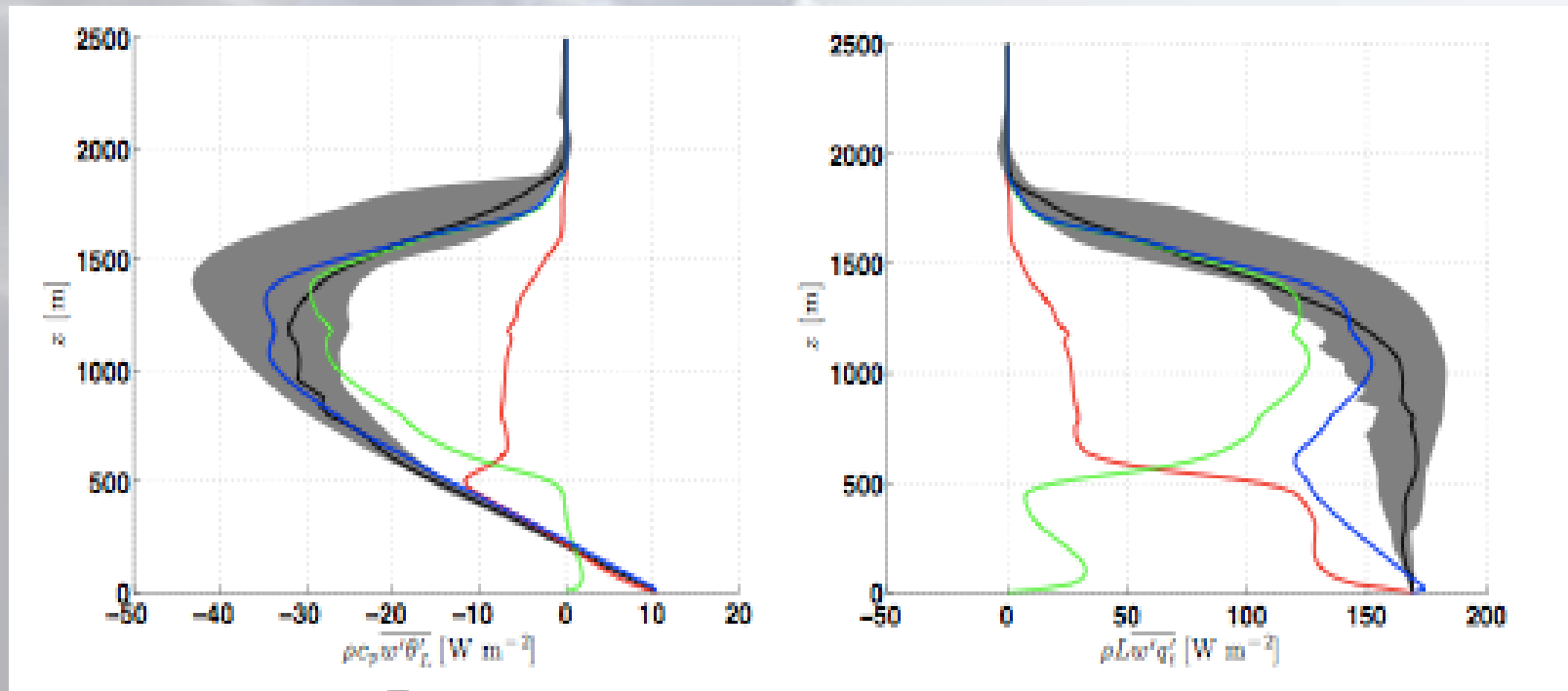


- 20 m (solid)
- 30 m (dashed)
- 40 m (dotted)
- 60 m (dash-dotted)

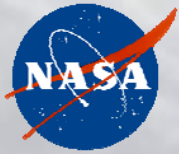


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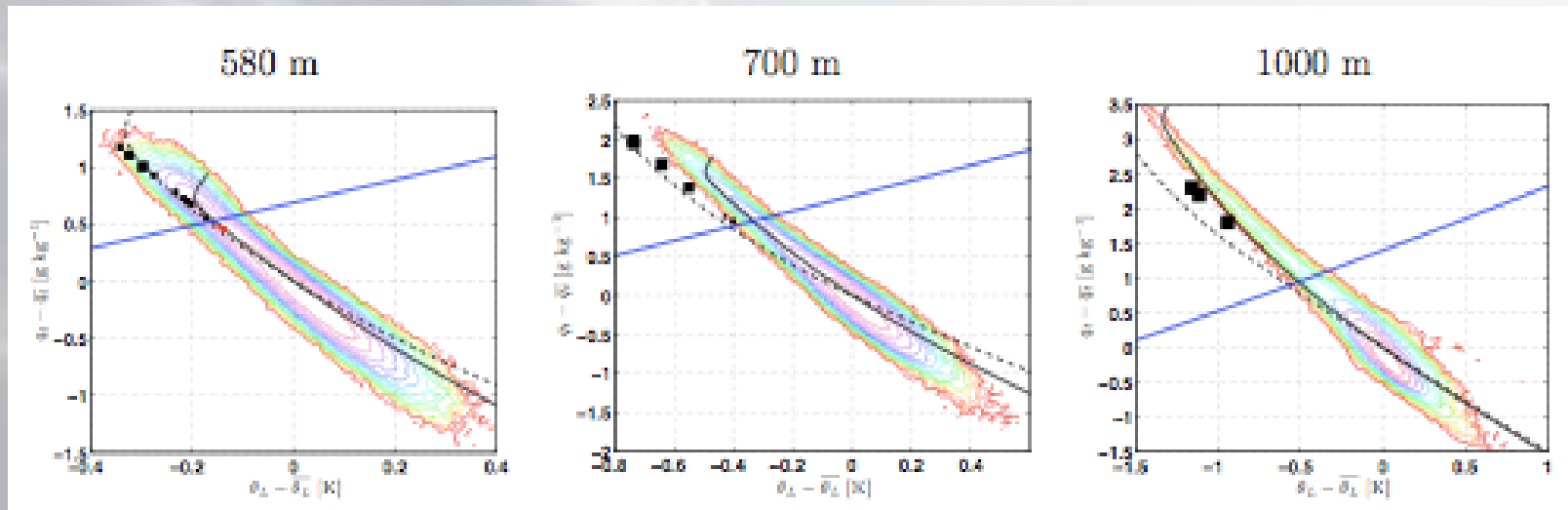
EDMF using PDF of updraft properties: BOMEX and the vertical fluxes



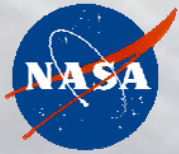
- LES (black)
- ED (red)
- MF (green)
- EDMF (blue)



EDMF using PDF of updraft properties: BOMEX and moist conserved variables PDFs



- LES (colored isolines)
- LES environmental profile (solid black line)
- updraft values (red squares represent dry, black squares moist updrafts)
- SCM environmental profile (black dashed line)
- saturation line (blue line).

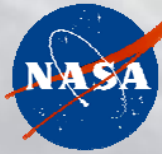


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A simple LES framework to study Sc , Cu and the transition

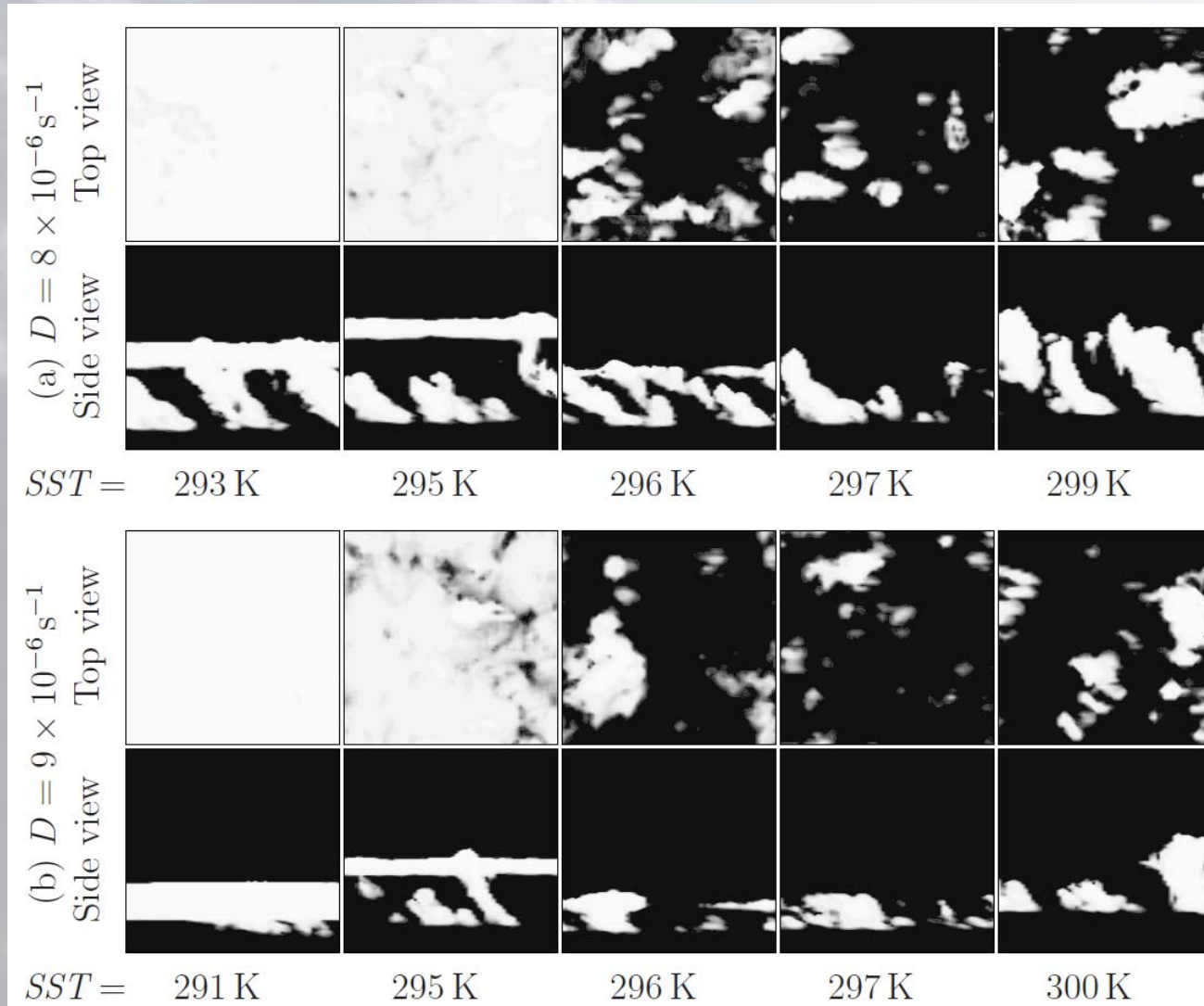
- ASTEX initial conditions (Duynderke et al. 1999)
- SST Monin—Obhukov surface boundary conditions
- Statistically steady (12 days)
- Imposed Large-scale advection and subsidence
- 2 K/day uniform clear-sky longwave cooling
- Cloud longwave cooling (Duynderke et al. 1999)
- 3.2 km x 3.2 km x 3 km domain
- 20 m x 50 m x 50 m resolution
- 10 cases: 5 SSTs, 2 Divs



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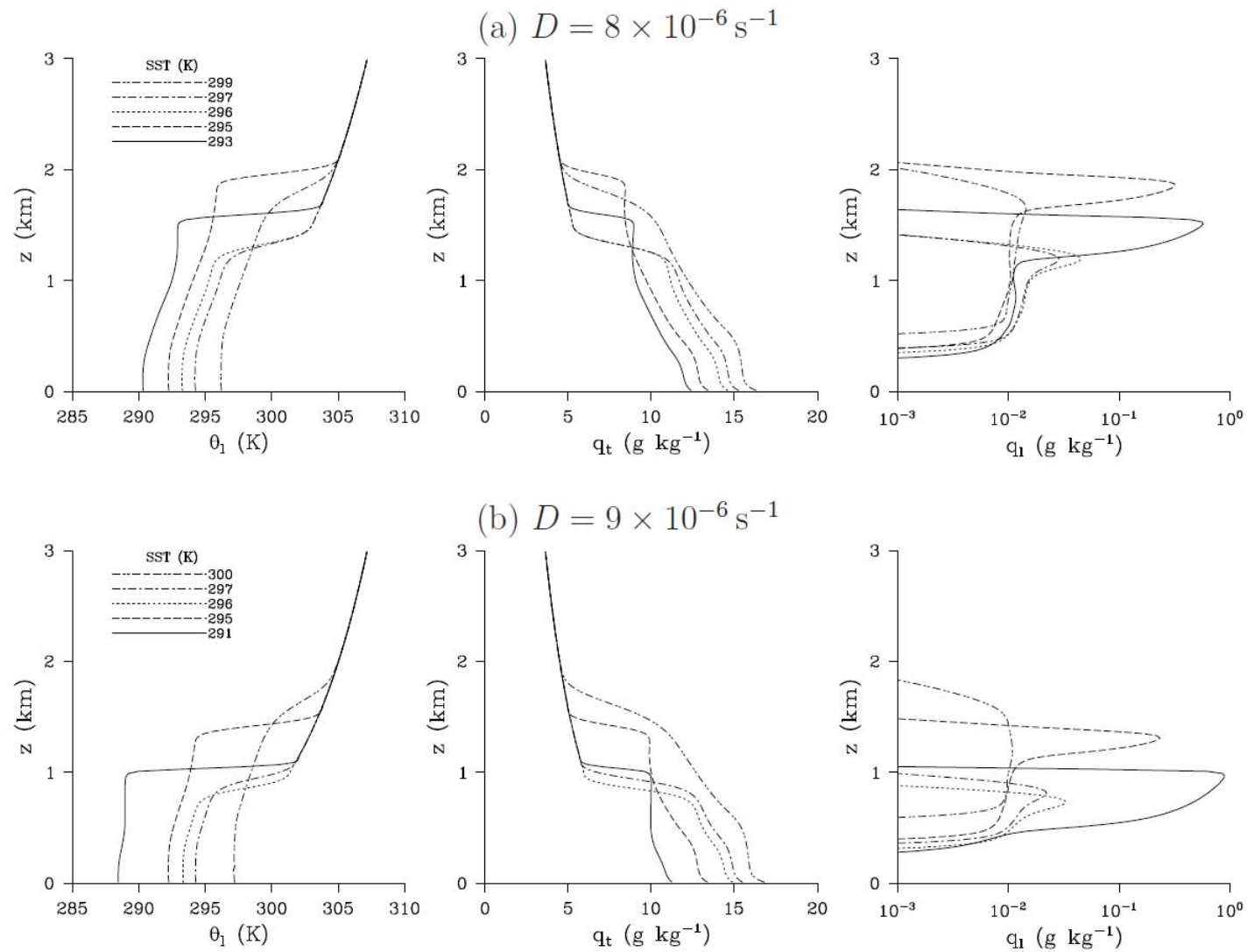
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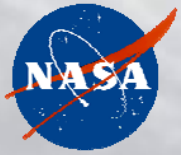
How does it look like?





Mean thermodynamic profiles

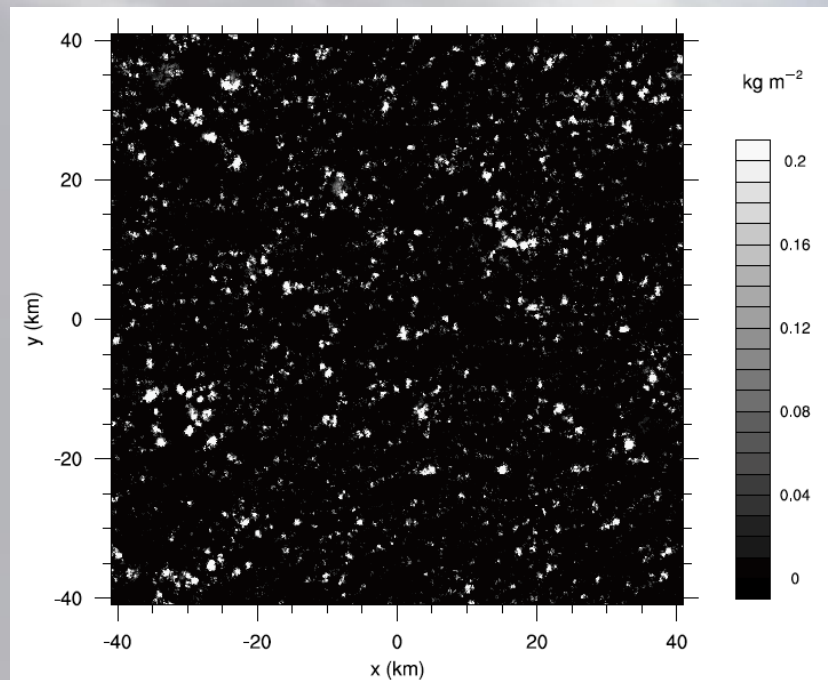




Large-domain LES: RICO

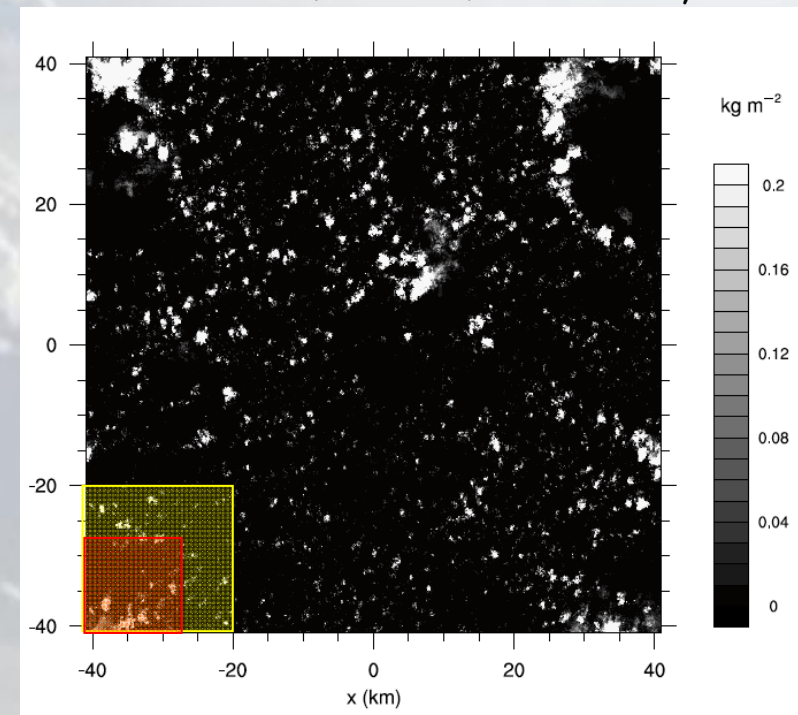
- Domain size $80 \times 80 \times 4$ km
- Resolution is 20 m, uniform
- $4096 \times 4096 \times 200 = 3.3$ billion cells

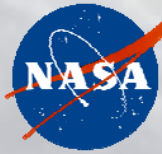
$t = 10$ h



$t = 15$ h

Matheou et al, 2011

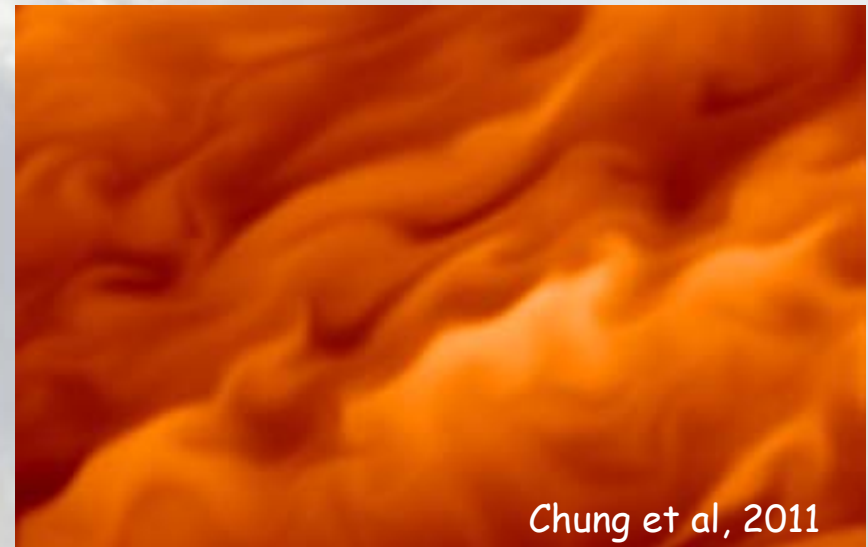
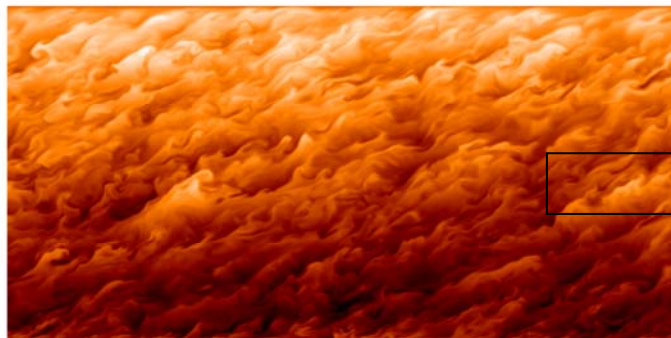
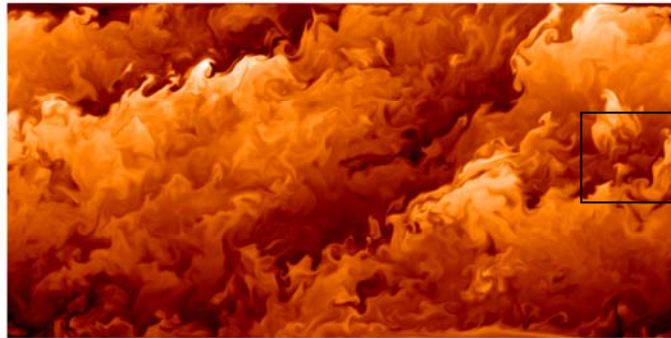




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Direct Numerical Simulation (DNS) of stratified homogeneous turbulence

Stratification/stability
increases

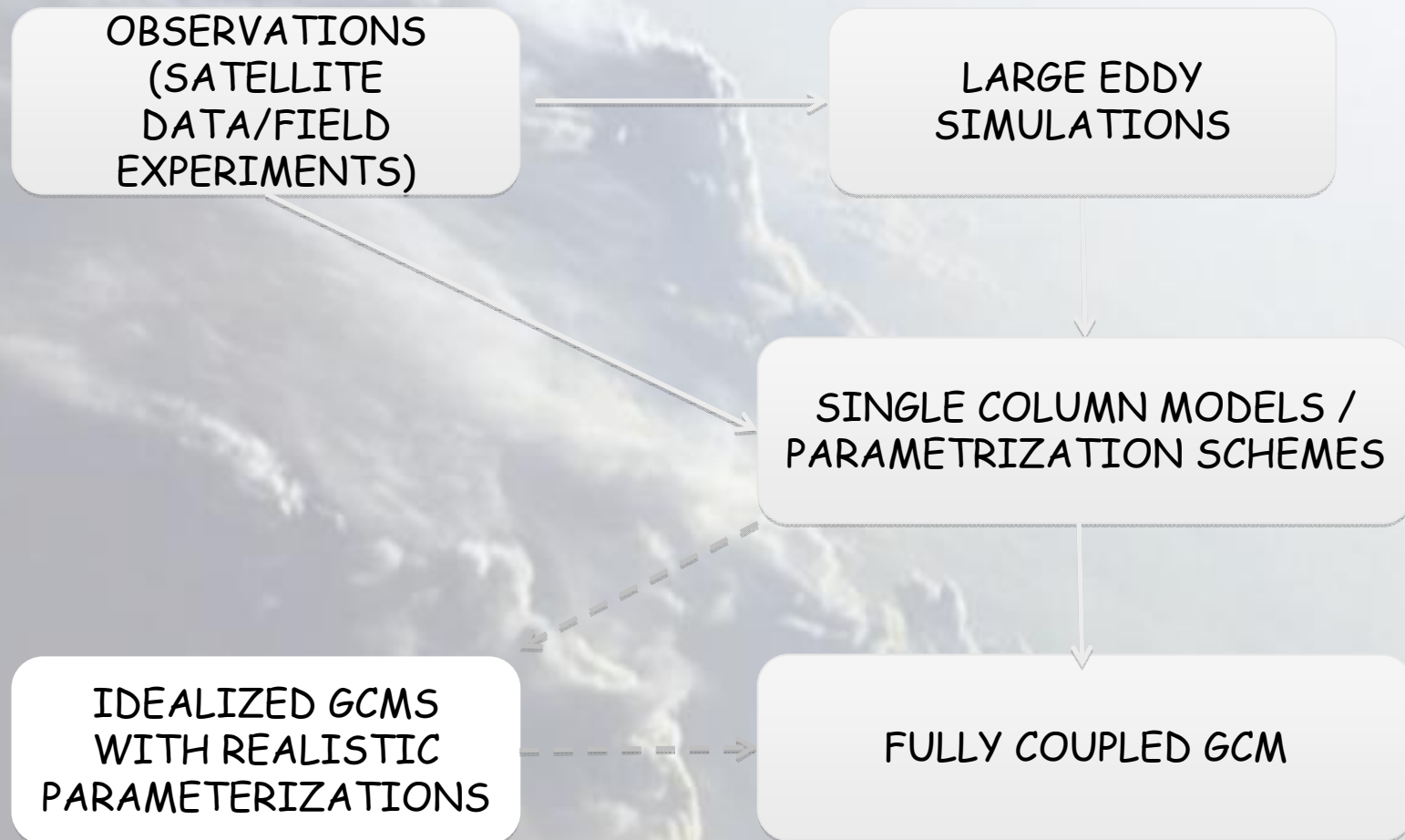


Chung et al, 2011



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Motivation for implementing EDMF in simple GCMs



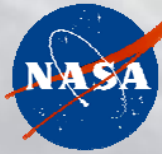


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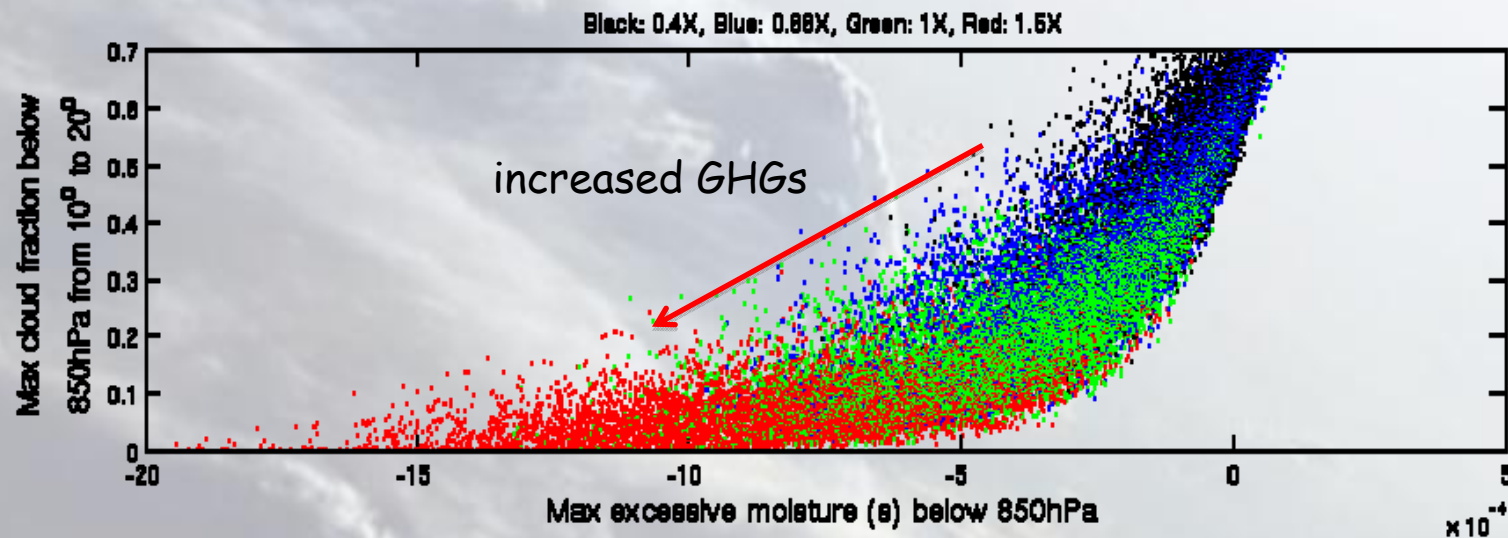
Simple GCM Description

- GFDL idealized GCM with a slab-ocean
- No diurnal or seasonal cycle
- No large-scale condensation
- No deep convection
- EDMF for sub-grid vertical mixing (dry, shallow and deep convection)
- Gaussian PDF-based cloud parameterization
- Vary the longwave optical depth (LWOD) to represent climate change with changing GHG

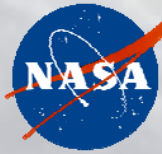


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Cloud changes for increased GHGs in simple GCM



- Subtropical low clouds decrease with with warming climate - increased GHGs
- Lower tropospheric stability does not appear related to this change
- Decrease of low cloud fraction is related to decrease of $s = q_t - q_s$



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Summary

- 1) A variety of projects related to EDMF development at JPL/Caltech
- 2) EDMF and TKE have been combined to represent dry and cumulus boundary layers
- 3) Using PDF of updraft properties leads to new EDMF shallow convection approaches: deterministic and stochastic sampling of cloud base PDF
- 4) LES steady-state simulations of Sc , Cu and transition help EDMF evaluation
- 5) EDMF and PDF-clouds have been implemented in idealized GCM to perform climate change investigations



Eddy-Diffusivity/Mass-Flux Model

Dividing a grid square in two regions (updraft and environment) and using Reynolds decomposition and averaging leads to

$$\overline{w'\phi'} = a_u \overline{w'\phi'_u} + (1-a_u) \overline{w'\phi'_e} + a_u(1-a_u)(w_u - w_e)(\phi_u - \phi_e)$$

where a_u is the updraft area. Assuming $a_u \ll 1$ and $w_e \sim 0$ leads to

$$\overline{w'\phi'} = \overline{w'\phi'_e} + a_u w_u (\phi_u - \bar{\phi})$$

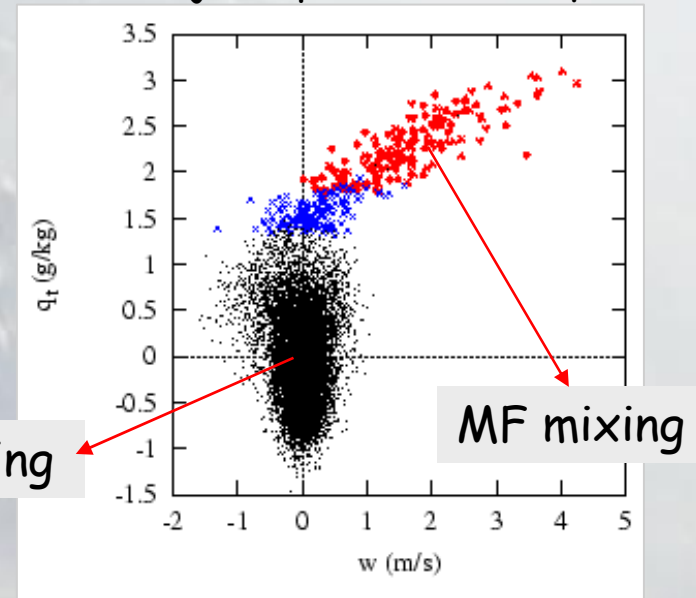
ED closure: assuming ED for 1st term and neglecting 2nd term

MF closure: neglecting 1st term and assuming $M = a_u w_u$

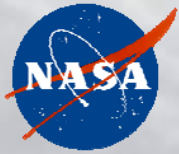
$$\text{EDMF: } \overline{w'\phi'} = -k \frac{\partial \bar{\phi}}{\partial z} + M(\phi_u - \bar{\phi})$$

Siebesma & Teixeira, 2000

Bimodal joint pdf of w and q_t



EDMF is able to reproduce variety of boundary layer convection types



Cloud and Convection Parameterization: Moist conserved variables

Traditional set of thermodynamic variables

$$\frac{\partial \theta}{\partial t} = -\frac{\partial}{\partial z}(\overline{w'\theta'}) + \frac{L}{C_p} \frac{\theta}{T} C \quad \frac{\partial q}{\partial t} = -\frac{\partial}{\partial z}(\overline{w'q'}) - C \quad \frac{\partial l}{\partial t} = -\frac{\partial}{\partial z}(\overline{w'l'}) + C$$

θ - potential temperature, q - specific humidity, l - liquid water

For convenience:
the mean of a
variable $\overline{\varphi}$ is often
represented as φ

Moist conserved variables

$$\frac{\partial \theta_l}{\partial t} = -\frac{\partial}{\partial z}(\overline{w'\theta_l'}) \quad \frac{\partial q_t}{\partial t} = -\frac{\partial}{\partial z}(\overline{w'q_t'})$$

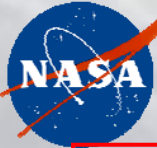
$$q_t = q + l$$

Total water content

$$\theta_l = \theta \left(1 - \frac{L}{C_p T} l \right) \quad \text{Liquid water potential temperature}$$

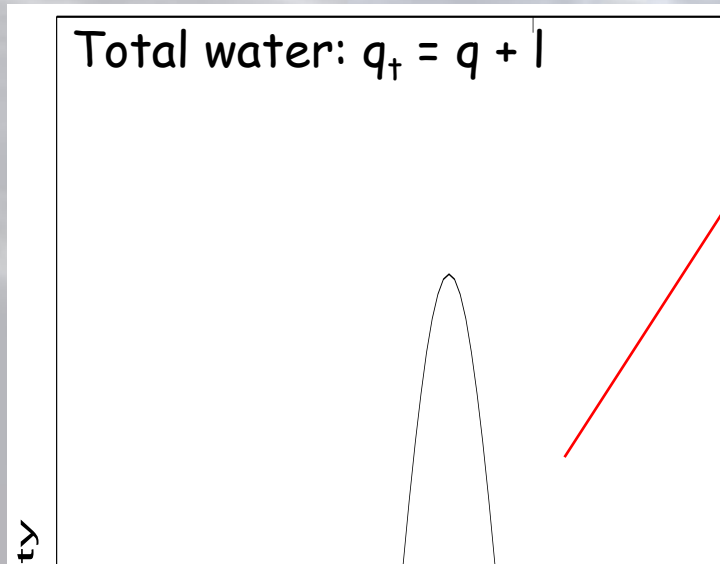
Two major advantages of using conserved variables:

- 1) The cloud/condensation term disappears from the equations
- 2) The ED approach is able to represent the correct cloud fluxes



PDF-based Cloud Parameterization

PDF cloud parameterizations are based on the pdf of q_t (in this simple example) or on the joint pdf of q_t and θ_l



Values larger than saturation are cloudy

$$a = \int_{q_s}^{+\infty} p(q_t) dq_t$$

$a = \text{cloud fraction}$

$$\bar{l} = \int_{q_s}^{+\infty} (q_t - \bar{q}_s) p(q_t) dq_t$$

Mellor, 77; Sommeria & Deardorff, 77

Gaussian PDF leads to cloud fraction and liquid water as a function of Q :

$$a = \frac{1}{2} + \frac{1}{2} \operatorname{erf} \left(\frac{Q}{\sqrt{2}} \right)$$

$$\frac{l}{\sigma} = aQ + \frac{1}{\sqrt{2\pi}} e^{-Q^2/2}$$

$$Q = \frac{q_t - q_s}{\sigma}$$



Pdf-based cloud parameterizations

How to determine the variance of total water?

1) Prognostic equation:

$$\frac{\partial}{\partial t} (\overline{q_t' q_t'}) = -2 \overline{w' q_t'} \frac{\partial q_t}{\partial z} - \frac{\partial}{\partial z} (\overline{w' q_t' q_t'}) - \frac{\overline{q_t' q_t'}}{\tau_q}$$

2) Diagnostic equation:

$$\overline{q_t' q_t'} = -2 \tau_q \overline{w' q_t'} \frac{\partial q_t}{\partial z}$$

Eddy-diffusivity

$$\overline{q_t' q_t'} = 2 \tau_q k \left(\frac{\partial q_t}{\partial z} \right)^2$$

Mass-flux

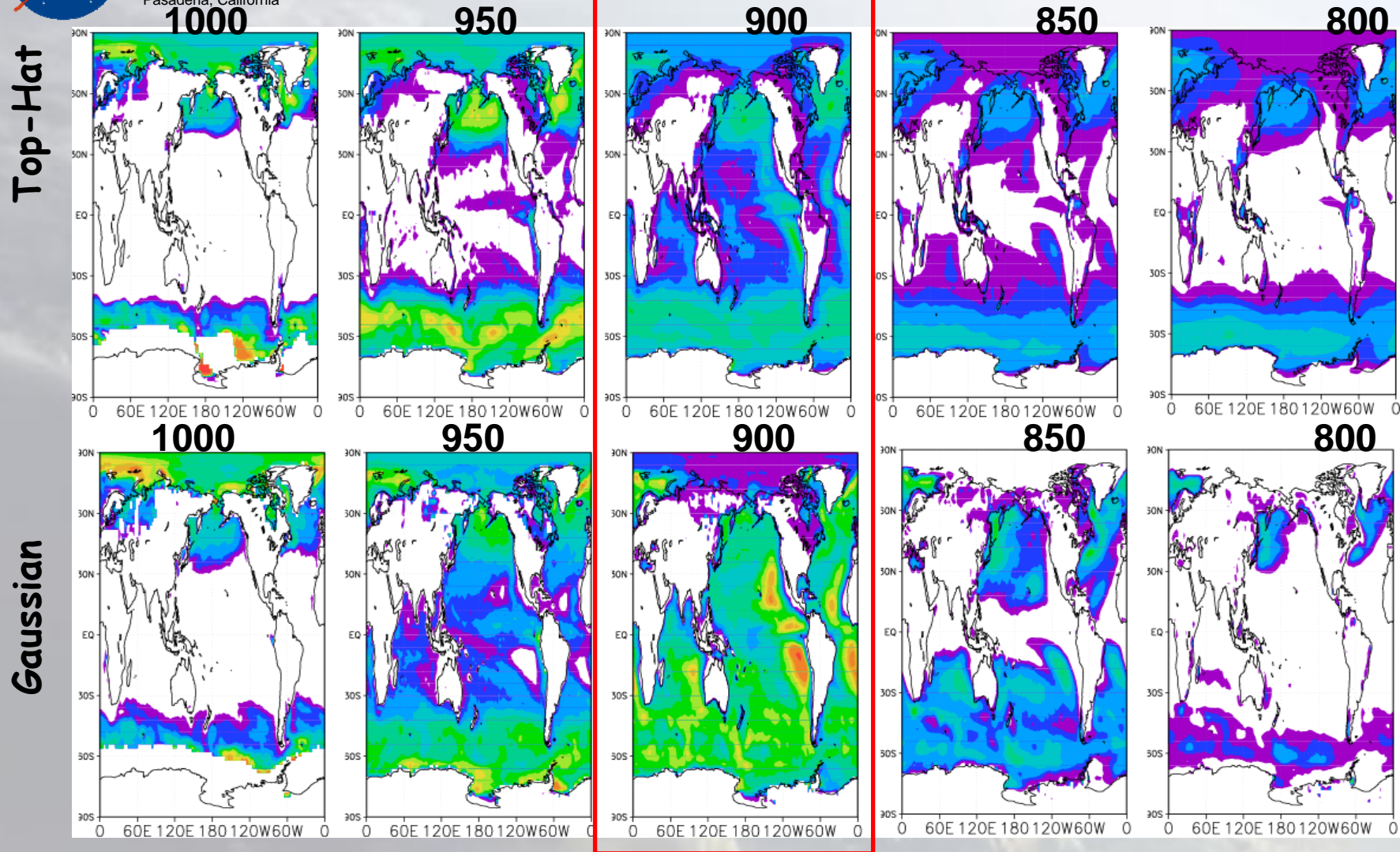
$$\overline{q_t' q_t'} = -2 \tau_q M (q_t^u - q_t) \frac{\partial q_t}{\partial z}$$



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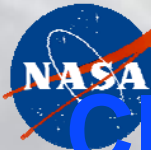
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Gaussian PDF scheme for PBL clouds implemented in GEOS-5: ED-based variance and interactive radiation



Two year simulation



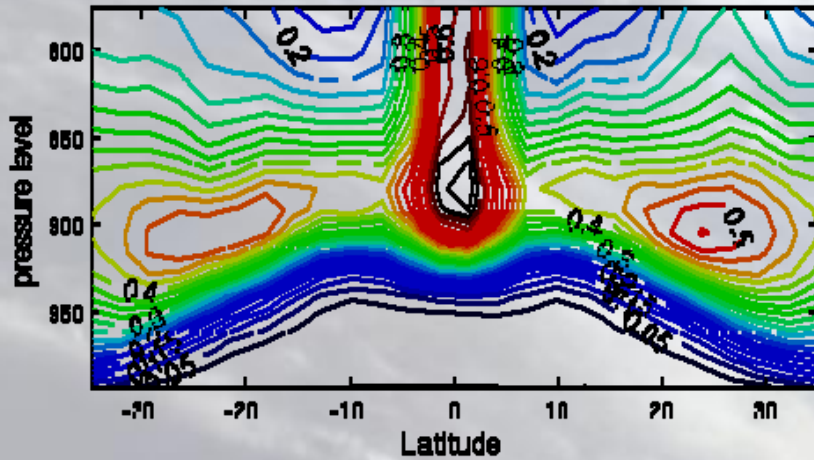


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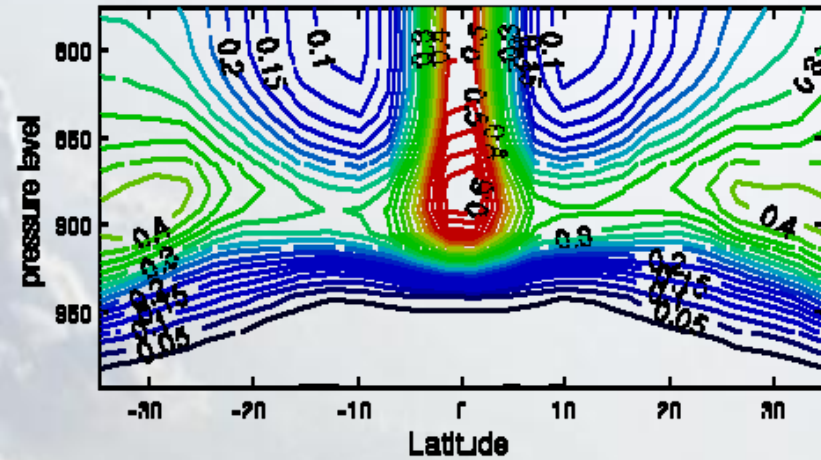
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Cloud-Fraction with different LWODs

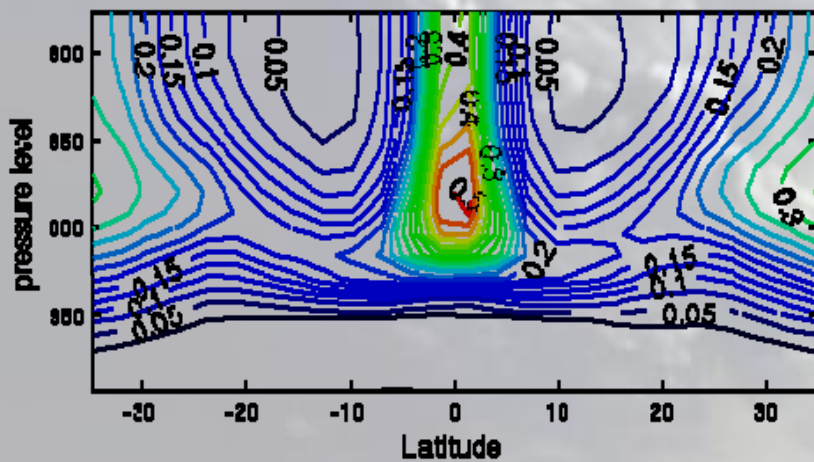
Cloud Fraction for 0.40x optical depth



Cloud Fraction for 0.66x optical depth



Cloud Fraction for 1.00x optical depth



Cloud Fraction for 1.50x optical depth

